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2020

Materna, K., Feng, L., Lindsey, E. O., Hill, E. M., Ahsan, A., Khorshed Alam, A. K. M., Kyaw, M. O., Than, O., Aung, T., Ngwe Khaing, S. & Bürgmann, R. (2020). GNSS characterization of hydrological loading in South and Southeast Asia. Geophysical Journal International, 224(3), 1742-1752. https://dx.doi.org/10.1093/gji/ggaa500

https://hdl.handle.net/10356/148918

https://doi.org/10.1093/gji/ggaa500

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Supplementary Material for: GNSS characterization of hydrological loading in South and Southeast Asia

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Text S1

For performing the synthetic calculation shown in Figure 6, we distributed ~14000 disk loads according to the locations of rivers and lakes in SE Asia that were identified using a 1-arcminute water mask; we then applied spatial loads of 6 km radius (matching our PREM Green's function for the rest of the computation) and height of 17 cm of water on each of those disks. The total imposed volume is 300 km^3, roughly equivalent to the excess discharge in the Ganges river during the monsoon season (Jian et al., 2009). This is intended as an order-of-magnitude calculation rather than an exact budget of the hydrological cycle.

Jian, J., Webster, P. J., & Hoyos, C. D. (2009). Large-scale controls on the Ganges and Brahaputra river discharge on intraseasonal and seasonal time-scales. *Quarterly Journal of the Royal Meterological Society*, 135, 353–370.

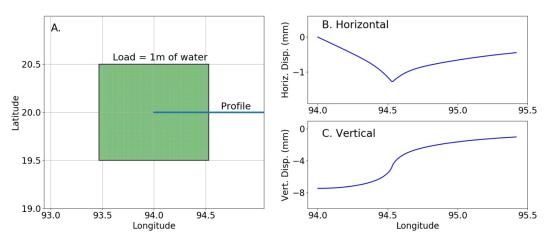


Figure S1: An example of a 1° x 1° loading calculation on a PREM earth structure. The maximum vertical deformation due to the load is several times larger than the maximum horizontal deformation.

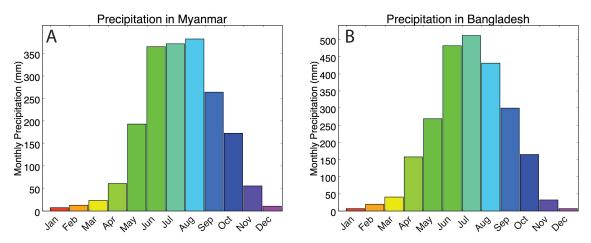


Figure S2: Average monthly precipitation in Myanmar and Bangladesh (World Bank Group Climate Change Knowledge Portal, 2010).

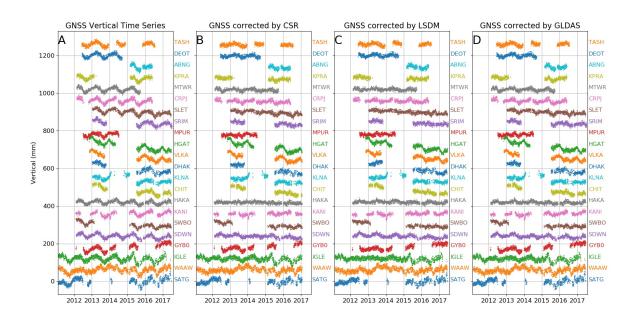


Figure S3: Vertical GNSS time series at the MIBB network, corrected for hydrological loading using each of the models explored in this paper.

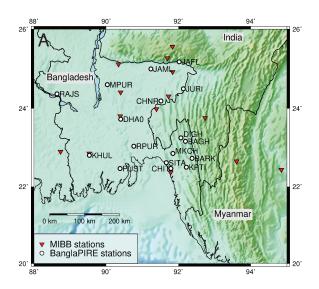


Figure \$4: Map of the BanglaPIRE GNSS network.

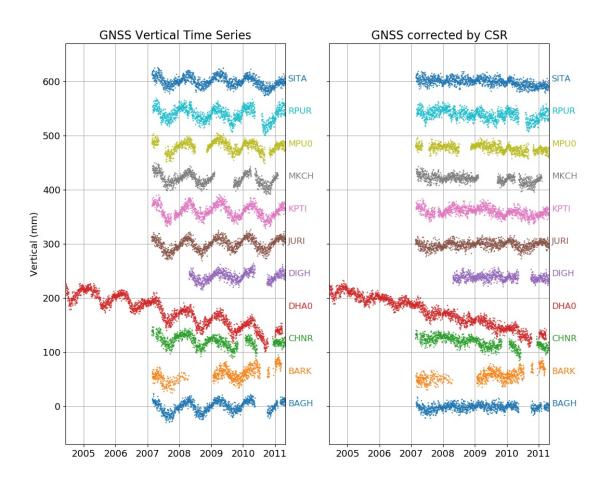


Figure S5: Vertical data at the BanglaPIRE GNSS stations showing the corrected GNSS time series after applying a GRACE-based hydrological loading model.

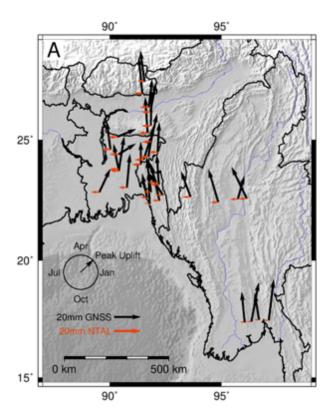


Figure S6: Phasor diagram for the raw GNSS time series before correction of non-tidal atmospheric and non-tidal ocean loading. The orange vectors are the phasors for the applied non-tidal ocean and atmosphere correction. The ocean and atmosphere loading is small and has peak uplift around July, which is consistent with a low-pressure system at the time of the monsoon.

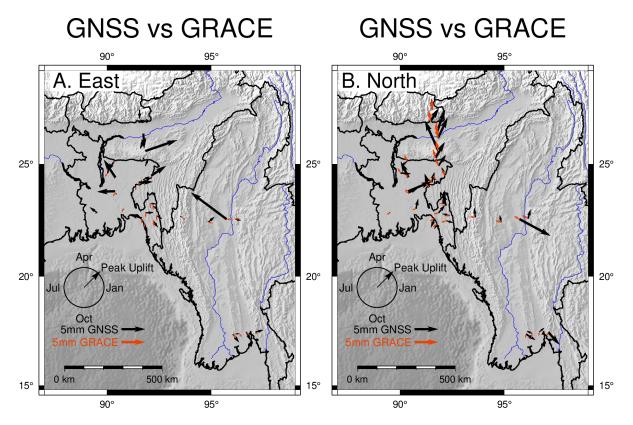


Figure S7: Estimated phase and amplitude of horizontal seasonal oscillations in the GNSS data (black) and the GRACE models (orange).

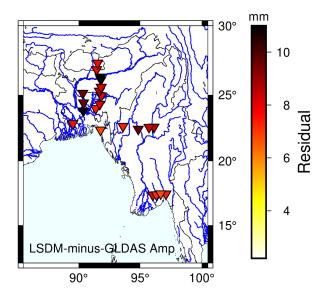


Figure S8: The amplitude difference between LSDM and GLDAS model predictions at each station in the MIBB network. The differences between these models is hypothesized to be the loading component from rivers and lakes. The scale bar is the same as Figure 6a.

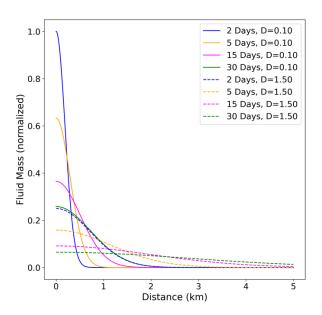


Figure S9: Diffusion curves for a sudden influx of surface water diffusing into the subsurface. D is the hydraulic diffusivity. The normalized flux on the y axis is unitless. Hydraulic diffusivities of this range are consistent with sandstone, basalt, chalk, and karst and inconsistent with granite, marble, and clay (Barbour and Wyatt, 2014). In a time period of 15-30 days, water could reach a subsurface aquifer at several kilometers' depth but could not move significant distances laterally.

| Station (Network) | Reason |
|-------------------|----------------------------|
| DRNG (MIBB) | Time series is mostly gaps |
| ICHA (MIBB) | Insufficient data length |
| BNGM (BanglaPIRE) | Insufficient data length |
| CHIO (BanglaPIRE) | Time series is mostly gaps |
| JAFL (BanglaPIRE) | Insufficient data length |
| JAML (BanglaPIRE) | Insufficient data length |
| KHUL (BanglaPIRE) | Time series is mostly gaps |
| PUST (BanglaPIRE) | Time series is mostly gaps |
| RAJS (BanglaPIRE) | Time series is mostly gaps |

Table S1: Blacklist of unused GNSS stations for hydrological loading analysis.

Table S2: Amplitude, phase, and velocity results for various models and the GNSS observations from MIBB and BanglaPIRE stations (included as separate text file).