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1 **Average velocity of solitary coarse grain in flows**  
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21   **Abstract.** A formula is derived for the prediction of the average velocity of a solitary coarse  
22   grain moving over a rigid flat channel bed. To verify the theoretical consideration, data  
23   collected in the present study and reported in the literature (811 sets in total) are analyzed for  
24   typical bed roughness configurations. The results show significant improvements in the  
25   prediction of solitary grain velocity for both smooth and rough bed conditions, in comparison  
26   with other formulas available in the literature. From the proposed formula, it follows that for  
27   rough bed conditions, the dimensionless grain velocity depends on the Shields number and  
28   also the ratio of bed roughness height to the grain diameter. The present study is limited to  
29   motion of solitary grains, but the result obtained implies that bedload transport rate could be

30 also affected by the equivalent bed roughness height even under mobile bed conditions and  
31 such effects may become pronounced for non-uniform sediments.

32 Keywords: grain velocity; bedload; smooth bed; rough bed; bed roughness; bed shear stress

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34

## 35 **Introduction**

36 Bedload transport rate can be estimated as a product of average grain velocity, bedload  
37 thickness and sediment concentration (Bagnold, 1973, Bridge and Dominic, 1984). However,  
38 how to quantify each individual factor, theoretically or even experimentally, still remains  
39 challenging at present because of complex near-bed flow phenomena and fluid-grain and  
40 grain-grain interactions. To simplify relevant physical phenomena, laboratory observations of  
41 sediment motion near the bed could focus only on a solitary bedload grain. Such  
42 simplifications do not result in a general formula for the computation of bedload transport  
43 rate, but they provide important insights into dynamics of the near-bed sediment motion  
44 (Francis, 1973).

45 In this paper, all information presented will be largely associated with a solitary grain  
46 moving near the bed. The grain velocity is defined as the average velocity of the grain  
47 moving continuously over a channel reach, of which the length is a few orders greater than  
48 the grain diameter. First, previous studies related to the calculation of bedload grain velocity  
49 are reviewed. Then, a new formula is derived by considering forces exerting on a moving  
50 grain in the context of spatial average. To substantiate the theoretical consideration,  
51 experimental data from the literature and collected in the present study are analyzed for  
52 typical bed configurations. The results show that the new formula represents well observed  
53 bedload grain velocities in flows over smooth and rough channel beds of different roughness.

54

55

## 56 **Previous Formulas**

57 Several empirical and analytical formulas for predicting bedload grain velocity have been  
58 reported in the literature, as summarized in Table 1. One of the early efforts is due to Einstein  
59 (1950), who proposed that the average bedload velocity is proportional to the shear velocity  
60 with respect to the grain. Bagnold (1973) assumed that the difference between the local fluid  
61 velocity at the representative height of a saltating grain and the grain velocity can be  
62 approximated to be the settling velocity. Fernandez Luque and van Beek (1976), Abbott and  
63 Francis (1977) and Bridge and Dominic (1984) developed their formulas individually based  
64 on the fact that the observed grain velocity appears to be linearly related to the bed shear  
65 velocity. Some regression analyses have been done by van Rijn (1984), Wu et al. (2006) [also  
66 see Wu (2008)], and Julien and Bounvilay (2013), and their formulas implicitly include  
67 viscous effects of fluid on the motion of bedload grain near the bed.

68 For comparison purposes, the previous formulas in Table 1 are presented in terms of the  
69 dimensionless grain velocity  $\chi$  defined as

$$\chi = \frac{u_s}{\sqrt{\Delta g D}} \quad (1)$$

70 and the dimensionless bed shear stress or Shields number  $\theta$ ,

$$\theta = \frac{u_*^2}{\Delta g D} \quad (2)$$

71 In Eqs. (1) and (2),  $u_s$  is the grain velocity,  $u_*$  is the shear velocity,  $\Delta [= (\rho_s - \rho)/\rho]$  is the  
72 effective specific density,  $\rho$  is the fluid density,  $\rho_s$  is the grain density,  $g$  is the gravitational  
73 acceleration and  $D$  is the grain diameter. The other dimensionless parameters used include  
74 the dimensionless grain diameter  $D_*$ ,

$$D_* = \left( \frac{\Delta g}{\nu^2} \right)^{1/3} D \quad (3)$$

75 where  $\nu$  is the kinematic viscosity of fluid, and transport stage parameter  $T_*$ ,

$$T_* = \frac{\theta - \theta_c}{\theta_c} \quad (4)$$

76 where  $\theta_c$  is the critical Shields number for the incipient sediment motion.

77 By comparing the previous formulas, two observations can be made. First, the critical bed  
 78 shear stress for the incipient sediment motion has been used in almost all the previous studies.  
 79 As a result, the formulas would fail to predict the grain velocity if the bed shear stress is less  
 80 than the critical value for initial sediment motion. The application of the concept of the  
 81 critical bed shear stress in the formulas is based on the argument that a bedload grain does not  
 82 move if the bed shear stress is less than its critical value. However, this argument may not  
 83 hold generally. After reviewing historical data for the incipient sediment motion, Lavelle and  
 84 Mofjeld (1987) concluded that the critical shear stress should not be included as an essential  
 85 parameter in the investigation of bedload transport rates. Paintal (1971) observed that there is  
 86 no shear stress below which no single grain moves. Laursen (1999) noticed that the same  
 87 sediment particle has as many values of the critical shear stress as there are sediment  
 88 transport formulas. Therefore, it is not necessary to involve the critical bed shear stress in the  
 89 formulation of the bedload grain velocity as well as the bedload transport rate (Cheng, 2002).

90 The second observation is that only Julien and Bounvilay (2013) included the effect of the  
 91 bed roughness on the bedload grain velocity, and the other formulas were basically developed  
 92 for the condition of the moving grain being the same as the sediment bed material. However,  
 93 it should be noted that the formula by Julien and Bounvilay (2013) is not applicable for  
 94 smooth beds with the bed roughness size approaching zero.

95 In this study, both smooth and rough bed configurations are considered in performing  
 96 laboratory experiments and data analysis, and also in developing a new formula for

97 calculating solitary grain velocity. In particular, the critical bed shear stress or velocity is not  
98 directly involved in the calibration of the formula proposed. With 811 sets of experimental  
99 data compiled from 11 sources, the evaluation of the bedload grain velocity is discussed in  
100 detail for different bed configurations.

101

102

### 103 **Theoretical Consideration**

104 First, the grain velocity under consideration is taken to be the average velocity of the grain  
105 that moves over a channel reach. Being different from the local grain velocity, the average  
106 grain velocity is one-dimensional and has no vertical and transverse components. Therefore,  
107 it can be described using a one-dimensional mathematical model. Here we consider a  
108 sediment grain of diameter  $D$  moving along the channel bed at average velocity  $u_s$ , which is  
109 usually subject to the drag force  $F_D$ , the lift force  $F_L$ , the bed resistance  $F_R$ , and the  
110 submerged weight  $W$  (see Fig. 1). The drag force is expressed as

$$F_D = C_D \frac{\pi D^2}{4} \frac{\rho(u_s - u_b)^2}{2} \quad (5)$$

111 where  $C_D$  is the drag coefficient and  $u_b$  is the unobstructed, near-bed flow velocity that is  
112 evaluated at a representative height of the moving grain. The submerged weight of the grain  
113 is

$$W = \Delta \rho g \frac{\pi D^3}{6} \quad (6)$$

114 The resistance  $F_R$  is due to the grain-bed interaction and can be expressed as

$$F_R = W \cos \beta \tan \alpha \quad (7)$$

115 where  $\beta$  is the slope angle,  $\tan\alpha$  is the friction coefficient and  $\alpha$  is the friction angle. If  $u_b > u_s$ ,  
 116  $F_D$  is in the direction towards downstream. Therefore, the following momentum equation  
 117 applies in the streamwise direction,

$$F_D + W \sin \beta = F_R \quad (8)$$

118 It is noted that the lift force is ignored here. Locally, the lift force is important in the  
 119 investigation of local movements, e.g. a particular grain trajectory. However, in the one-  
 120 dimensional model, it is assumed that the grain moves only in the streamwise direction. As a  
 121 result, all forces are taken to be space-averaged in the channel reach under consideration.  
 122 This implies that the bed-normal lift force can be excluded for consideration.

123 Substituting Eqs. (5) to (7) into Eq. (8),

$$\chi = \frac{u_b}{\sqrt{\Delta g D}} - \sqrt{\frac{4 \cos \beta \tan \alpha - \sin \beta}{3 C_D}} \quad (9)$$

124 Eq. (9) applies for the condition of  $u_b > u_s$  and the second term on the right-hand-side implies  
 125 that the slope angle is smaller than the friction angle, i.e.  $\beta < \alpha$ . By taking  $u_b = c_1 u_*$ , where  
 126  $c_1$  is a coefficient and assuming that  $\beta$  is much smaller than  $\alpha$ , Eq.(9) can be rewritten as

$$\chi = c_1 \sqrt{\theta} - \sqrt{\frac{4 \tan \alpha}{3 C_D}} \quad (10)$$

127 It is noted that equations similar to Eq. (9) or (10) have been also proposed previously by van  
 128 Rijn (1984), Bridge and Dominic (1984) and Fredsoe and Deigaard (1992). Both Eqs. (9) and  
 129 (10) are simple but not practically useful. This is because difficulties are often encountered in  
 130 the evaluation of the three coefficients,  $c_1$ ,  $\tan\alpha$  and  $C_D$  due to complex dynamics of near-bed  
 131 flow and grain-bed interaction. In the following, we will first discuss to what extent the  
 132 coefficients can be fixed, and then simplify Eq. (10) for some cases.

133

134 **Estimate of  $c_1$**



135 By noting that  $c_1$  is the ratio of  $u_b$  to  $u_*$ , its value depends on the choice of a representative  
 136 distance  $y_b$  (from the center of the moving grain to the channel bed), at which  $u_b$  is  
 137 determined. For a flat bed, whether smooth or rough,  $c_1$  can be estimated according to the  
 138 near-bed velocity profile (Cheng, 2008),

$$c_1 = \left\{ (y_b^+)^{-3} + \left[ 2.5 \ln \left( 1 + \frac{9y_b^+}{1 + 0.3k_s^+ (1 - e^{-0.04k_s^+})} \right) \right]^{-3} \right\}^{-1/3} \quad (11)$$

139 where  $y_b^+ = u_* y_b / \nu$ ,  $k_s^+ = u_* k_s / \nu$  and  $k_s$  is the equivalent roughness height of the channel bed.  
 140 Eq. (11) describes the linear velocity distribution in the viscous sublayer, the logarithmic  
 141 profile in the inertial flow zone, and also the transition in between. From Eq. (11), it follows  
 142 that the calculation of  $c_1$  is subject to the representative height  $y_b$ , which varies in the order of  
 143  $D$ , and equivalent roughness height  $k_s$ , which depends on the bed material and its mobility.

144

#### 145 **Variation of $\alpha$ with bed roughness**

146 For natural sediments, the friction angle is markedly affected by particle shape and size  
 147 distribution. Wiberg and Smith (1987) suggested that the friction angle varies between  $50^\circ$   
 148 and  $60^\circ$  for well-sorted natural sands. Johnston et al. (1998) conducted in situ determination  
 149 of friction angles of fluvial gravels, showing that the average friction angle varied from  $49.1^\circ$   
 150 to  $61.9^\circ$  for median sediment grains. Similar results have been also reported by Buffington et  
 151 al. (1992). Significant changes in the friction angle occur when the grain in motion is  
 152 different from the bed sediment. A few theoretical and empirical formulas have been  
 153 proposed in previous studies to relate  $\alpha$  to the ratio of  $k_s$  to  $D$ . For example, by considering a  
 154 sphere of diameter  $D$  moving from its position of rest in a bed of spheres of diameter  $k_s$ ,  
 155 Miller and Byrne (1966) showed that the friction coefficient can be theoretically related to  
 156  $D/k_s$ ,

$$\tan \alpha = \frac{0.866}{\sqrt{(D/k_s)^2 + 2(D/k_s) - 0.333}} \quad (12)$$

157 Eq. (12) can be further approximated in terms of the power-law. For example, for  $k_s/D < 1$ ,

158 Eq. (12) is close to the following function,

$$\tan \alpha = 0.54 \left( \frac{k_s}{D} \right)^{0.8} \quad (13)$$

159 or

$$\alpha = 28 \left( \frac{k_s}{D} \right)^{0.75} \quad (14)$$

160 Experimental data show that Eqs. (12) and (13) generally underestimate the friction angle but  
 161 predict the trend of its variation with  $k_s/D$  reasonably (Miller and Byrne, 1966). The power-  
 162 law function, with varying coefficient and exponent, has been found very useful in  
 163 representing field measurements of the friction angle (Buffington, Dietrich and Kirchner,  
 164 1992). In the subsequent analysis, the power-law relation between  $\tan \alpha$  and  $k_s/D$ , with  
 165 undefined coefficient and exponent, will be used to simplify Eq. (10).

166

### 167 **Estimate of $C_D$**

168 For the motion of a grain without any boundary effect, the drag coefficient  $C_D$  can be  
 169 calculated using empirical formulas (Cheng, 1997, Cheng, 2009). In the presence of a  
 170 boundary,  $C_D$  may increase with the grain proximity to the boundary, which is not well  
 171 understood at present. With the information of boundary effects on settling velocity  
 172 summarized by Chien and Wan (1999), the drag coefficient in the presence of the boundary  
 173 could be formulated as  $[1+(9/32)(D/s)]^2 C_D$ , where  $s$  is the distance from the center of the  
 174 moving grain to the boundary. With this simple function, it can be estimated that  $C_D$   
 175 increases by 60% if  $s/D = 1$  and by 20% if  $s/D = 3$ . In comparison, only a small increase in

176  $C_D$  was observed in laboratory measurements. For example, Nakagawa and Tsujimoto (1981)  
 177 reported that for a sphere placed on a rough bed composed of spheres of the same kind, the  
 178 value of  $C_D$  remained unchanged or increased slightly, in comparison with that for a free  
 179 falling sphere. Similar findings were also presented by Schmeeckle et al. (2007).

180 From the foregoing discussions, it follows that there are no complete approaches available  
 181 at present for determining  $c_1$ ,  $\tan\alpha$  and  $C_D$ . This renders that Eq. (10), though simple, is  
 182 inapplicable for predicting the bedload grain velocity directly. However, on the other hand,  
 183 Eq. (10) can be simplified for some cases. For example, for a smooth channel bed, for which  
 184  $k_s$  is much smaller than  $D$  and  $\tan\alpha$  approaches zero, Eq. (10) reduces to

$$\chi = c_1 \sqrt{\theta} \quad (15)$$

185 Another simplification can be made possible for the case of coarse bedload grains, e.g.  
 186 natural gravels of  $D \geq 2$  mm, for which the viscous effect on the drag coefficient can be  
 187 ignored and then  $C_D$  can be approximated as a constant (Cheng, 1997). If  $\tan\alpha$  is expressed as  
 188 a power-law function of  $k_s/D$ , similar to Eq. (13), Eq. (10) is rewritten as

$$\chi = c_1 \sqrt{\theta} - c_2 \left( \frac{k_s}{D} \right)^m \quad (16)$$

189 where  $c_2$  is a coefficient and  $m$  is an exponent, both being positive. It is noted that Eq. (16)  
 190 reduces to Eq. (15) when  $k_s$  is much smaller than  $D$ . As shown in the following section, most  
 191 of the experimental data compiled in the present study were collected either under smooth-  
 192 bed conditions or with coarse grains, and thus they can be used to evaluate  $c_1$ ,  $c_2$  and  $m$  in Eq.  
 193 (16).

194

195

## 196 **Data Analysis**

197 Experimental data used for analysis consist of two parts. The first part, comprising 96 sets,  
198 was generated from the experiments conducted in the present study, as described  
199 subsequently. The second part, comprising 715 sets, was reproduced from the previous  
200 studies (Abbott and Francis, 1977, Fernandez Luque and van Beek, 1976, Francis, 1973, Hu  
201 and Hui, 1996, Julien and Bounvilay, 2013, Lee, et al., 2000, Lee and Hsu, 1994, Niño and  
202 García, 1998, Niño, et al., 1994, Parsons, 1972). A summary of the data is shown in Table 2.

203 In the previous studies, most of the experiments were conducted under the condition of the  
204 moving grain taken the same as uniform sediment bed materials (Abbott and Francis, 1977,  
205 Francis, 1973, Lee, Chen, You and Lin, 2000, Lee and Hsu, 1994, Niño, García and Ayala,  
206 1994). To facilitate experimental observations, the sediment bed is often prepared by gluing a  
207 layer of sediment grains onto the channel bottom, which results in  $k_s \approx D$ . In comparison, the  
208 measurement of the grain velocity under the mobile bed condition is more complicated, and  
209 limited efforts have been made by Fernandez Luque and van Beek (1976) and Niño et al.  
210 (1994). Different from the others, Julien and Bounvilay (2013) and Bounvilay (2003)  
211 investigated effects of the bed roughness height on the grain velocity. They employed a series  
212 of moving grains of different sizes when the bed condition remained unchanged, which  
213 yielded the relative roughness height  $k_s/D$  varying from 0 to 1.0.

214 All the data can be categorized into four different cases according to the bed conditions  
215 (Fig. 2). The first case is related to the experiments conducted under the smooth bed  
216 condition, of which the bed roughness height  $k_s$  is much smaller than the grain diameter  $D$ , i.e.  
217  $k_s \ll D$ . The second case involves the investigations with the moving grain and bed material  
218 of the same size, i.e.  $k_s \approx D$ . The third case is limited to the case of  $k_s < D$ . In the first three  
219 cases, the bed, whether smooth or rough, is rigid. The last case concerns the experiments  
220 with mobile beds for which  $k_s > D$ .

221 In the following, individual data analysis is performed for each case to check applicability  
 222 of the foregoing theoretical consideration.

223

224 **Case 1: Smooth bed ( $k_s \ll D$ )**

225 The first case concerns a smooth channel bed, of which the bed roughness height is much  
 226 smaller than the grain diameter and thus the bed resistance is negligible. Four series of data  
 227 were selected under this condition.

228 The first series is due to Parsons (1972) who performed experiments with glass beads and  
 229 sand grains in laminar flows. To avoid any significant bed friction, only the data with glass  
 230 beads are used here. It should be mentioned that this is the only data series related to laminar  
 231 flows that is used in the present study. For laminar flows, the velocity profile near the bed is  
 232 linear and Eq. (11) reduces to

$$c_1 = \frac{u_* y_b}{\nu} \quad (17)$$

233 Furthermore, by noting that  $y_b$  is proportional to  $D$ , Eq. (10) can be rewritten as

$$\chi = c_3 \text{Re}_* \sqrt{\theta} - \sqrt{\frac{4 \tan \alpha}{3 C_D}} \quad (18)$$

234 where  $\text{Re}_* = u_* D / \nu$  and  $c_3$  is a constant. In Fig. 3,  $\chi$  is plotted against  $\text{Re}_* \sqrt{\theta}$  with the data  
 235 reported by Parsons (1972) for six different sizes of glass beads ( $D = 0.25\text{-}0.85$  mm). It  
 236 shows that the relationship between  $\chi$  and  $\text{Re}_* \sqrt{\theta}$  is almost linear, which is consistent with  
 237 Eq. (18). The results also imply that  $c_3 = 0.27$  (and thus  $y_b = 0.27D$ ) and the bed resistance,  
 238 represented by the last term on the RHS of Eq. (18), is relatively small.

239 Next, smooth-bed experiments performed under turbulent flow conditions are examined.  
 240 For turbulent flows, the ratio of  $u_b/u_*$  varies with  $y_b$  and  $k_s$  [see Eq. (11)]. It can be assumed  
 241 that  $y_b$  is proportional to  $D$  and an empirical  $k_s$ -value can be assigned to a particular smooth

242 bed. Computations using Eq. (11) show that  $u_b/u_*$  increases with increasing  $Re_*$ , but varying  
243 in a limited range for large  $Re_*$ . For example, if taking  $y_b = 0.27D$ , as suggested by the  
244 laminar flow result, and  $k_s = (1\%-10\%)D$ ,  $u_b/u_*$  would vary from 11 to 18 for  $Re_* > 60$ .  
245 However, it is not clear how to determine  $y_b$  generally for turbulent flows (Bagnold, 1973).

246 Three series of data for turbulent flows over a smooth bed are employed for analysis,  
247 which were generated from the experiments by Hu and Hui (1996) and Julien and Bounvilay  
248 (2013) and also by the present study. In Julien and Bounvilay's (2013) experiments (also see  
249 Bounvilay (2003)), tested grains comprised glass beads, steel balls and natural sands, but only  
250 the data for the glass beads are used here. This selection was made for two reasons. First, the  
251 experiments by Julien and Bounvilay's (2013) were conducted under relatively low bed shear  
252 stress conditions ( $\theta = 0.000076-0.016$ ), at which the grain velocity could be sensitive to  
253 channel bed resistance. Second, in comparison to glass beads, both natural sands and steel  
254 balls may induce significant bed resistance because of their angular shape or high density. In  
255 comparison, Hu and Hui (1996) carried out their observations with much higher bed shear  
256 stresses ( $\theta = 0.15-1.41$ ), for which the effect of the bed resistance on the grain velocity could  
257 be considered insignificant in spite of some non-spherical grains tested.

258 To obtain additional data in the range of the intermediate Shields number (say,  $\theta = 0.01-$   
259  $0.1$ ), similar smooth-bed experiments were also carried out in the present study in an open  
260 channel. The channel was 100 mm in width, 150 mm in height and 3000 mm in length, with  
261 both the channel wall and bed made of glass. Water was circulated using a centrifugal pump,  
262 providing a discharge in the range of 0-1.3 l/s, and the flow rate was measured with an  
263 electromagnetic flow meter. Water temperature was measured with a mercury thermometer of  
264 an accuracy of  $0.1^\circ\text{C}$  and the kinematic viscosity was calculated using  
265  $\nu = [60/(T + 40)]^{1.45} 10^{-6}$  (Cheng, et al., 2011), where T is in degree Celsius and  $\nu$  in  $\text{m}^2/\text{s}$ .

266 Three kinds of spheres with  $\rho_s/\rho$  varying from 1.14 to 2.60 were used, which all measured 11

267 mm in diameter. The flow depth was measured along the channel with a 0.25m interval by  
268 applying a pressure transducer through the sidewall. All the experiments were conducted  
269 under uniform flow conditions. To achieve a uniform flow, a channel slope was selected first  
270 after certain trial tests. Then the flow rate and tailgate were adjusted. This was a trial-and-  
271 error process and the slope, tailgate and flow rate could be adjusted several times for a  
272 particular test. The flow was assumed fully developed at the test section. This is because the  
273 flow depth was usually small. When  $h$  varied from 0.0125 m to 0.022 m (see the data  
274 included in Table 3), the distance from the beginning of the channel to the test section was  
275 equal to  $(45-80)h$  ( $= 1$  m). This agrees with the study by Ranga Raju et al (2000), who found  
276 that the developing length (required for 99% development of free stream velocity) varied  
277 from  $40h$  to  $120h$  for a range of aspect ratios and relative roughness heights.

278 For each test, a grain was first released near the bed at the beginning of the channel. To  
279 obtain the average grain velocity, the time duration was recorded for a grain to complete its  
280 journey over a 1.5m channel reach, which started from a section 1m downstream from the  
281 flume entrance. Each test was repeated 20 times under the same flow condition. The data  
282 collected in the present study are provided in Table 3.

283 It should be mentioned that sidewall effects were not investigated in this study. During the  
284 experiment, grains moving close to sidewalls were not used for measurements, and the  
285 relevant tests were repeated. It was observed that typically grains moved almost along the  
286 centerline of the channel, even under rough bed conditions. On the other hand, it is also noted  
287 that the grain was considered large in comparison with the flow depth. This may cause the  
288 experimental results different from those obtained under high submergence conditions. A  
289 further discussion in this respect will be provided subsequently.

290 The three series of data for turbulent flows over smooth beds are plotted in Fig. 4. It shows  
291 that the dimensionless grain velocity  $\chi$  can be linearly related to the square root of the Shields

292 number  $\theta$ . By ignoring the bed-resistance, the linear relationship can be described with Eq.  
293 (15) by taking  $c_1 = 13.1$  with a variation of about  $\pm 30\%$ , or in the range of 9.1 to 17.0. This  
294 range of  $c_1$  is close to that estimated with Eq. (11) for large  $Re^*$ . In processing the data, the  
295 shear velocity is taken as  $(ghS)^{0.5}$ , where  $h$  is the flow depth and  $S$  is the channel slope, by  
296 noting the large aspect ratios applied in all the three experiments.

297

## 298 **Case 2: Rigid sediment bed ( $k_s \approx D$ )**

299 In the second case, the size of the fixed bed roughness element is the same as that of the  
300 moving grain tested. Here a further differentiation is made between spherical and non-  
301 spherical grains. For spherical tested grain and bed material, 44 sets of data were collected in  
302 the present study (see Table 3). The experimental setup was the same as that for the smooth-  
303 bed tests but with a sediment bed, which was prepared with a layer of spheres ( $k_s = 11$  mm)  
304 placed with the densest arrangement over the entire flume bottom. Fig. 5 shows that for  
305 spherical grains, the measured grain velocity can be well represented using Eq. (16) with  $c_1 =$   
306 11.2 and  $c_2 = 1.0$ .

307 For non-spherical tested grain and bed material, in total 125 sets of data reported  
308 previously by Francis (1973), Abbott and Francis (1977), Lee and Hsu (1994), Niño and  
309 García (1998) and Lee et al. (2000) are used for analysis. Francis (1973) conducted six series  
310 of experiments using both natural and artificial grains of different shape and density. All the  
311 data are provided in Francis'(1973) paper but without information about the grain densities,  
312 which make inconvenient any further analysis of his measurements. However, for series K in  
313 his data, it is possible to estimate the grain densities from the measured settling velocity. This  
314 is explained as follows. The grains used in series K were cast in the shape of water-worn  
315 gravels, having the same size ( $D = 7.5$  mm) but different densities. By noting that the  
316 Reynolds number based on the settling velocity varied from 465 to 2798 for this series, it is



317 reasonable to take the drag coefficient to be 1.0 (Cheng, 1997) and then the densities can be  
318 estimated from the measured settling velocity. Therefore, the data of series K were used for  
319 the present analysis. As shown later, the dimensionless grain velocity so obtained appears to  
320 have similar variations to those reported in Abbott and Francis (1977), who continued the  
321 effort by Francis (1973). In addition, by considering availability of the necessary information  
322 about flow and grain properties, the data of set 3 (with  $D = 8.82$  mm,  $\rho_s/\rho = 1.2-2.6$ ) reported  
323 by Abbott and Francis (1977) were also reproduced from their figure for the present study.

324 All the experimental data for non-spherical grains with  $k_s \approx D$  can be fitted to Eq. (16)  
325 with  $c_1 = 12.5$  and  $c_2 = 1.5$ , as shown in Fig. 6. It is noted that this  $c_1$ -value is close to the  
326 average value ( $c_1 = 13.1$ ) obtained for turbulent flows in case 1. In comparison to  $c_1 (= 12.5)$   
327 for non-spherical grains, the value of  $c_1 (= 11.2)$  for spherical grains appears 10% smaller.  
328 This difference is largely caused by the grain shape, but it may be also associated with the  
329 small flow depth used in the present experiments with spherical grains. When the flow depth  
330 varies in the same order of the grain diameter, the channel resistance can increase  
331 significantly (Bathurst, et al., 1981, Ferguson, 2007). This implies that the grain velocity may  
332 reduce with a decrease in the ratio of the flow depth to grain diameter.

333

### 334 **Case 3: Rigid sediment bed ( $k_s < D$ )**

335 In the third case, the bed roughness height is smaller than the diameter of the tested grain.  
336 This case has been systematically studied by Julien and Bounvilay (2013) by varying  $k_s/D$   
337 between 0 and 1. Some of their data series for natural sediment grains are plotted in Fig. 7,  
338 showing variations of  $\chi$  with  $\theta$  at different  $k_s/D$ -values. Also superimposed in Fig. 7 are two  
339 straight lines, one with  $c_1 = 13.1$  for the case of  $k_s/D \approx 0$  and the other with  $c_1 = 12.5$  and  $c_2 =$   
340  $1.5$  for  $k_s/D \approx 1$ . It shows that the gradient of each line is almost close to each other but the  
341 intercept at the vertical axis varies with  $k_s/D$ . This implies that  $c_1$  in Eq. (16) can be

342 approximated to be a constant, e.g,  $c_1 \approx 13.0$ . Then, the values of  $c_2$  and  $m$  can be optimized  
343 based on the results shown in Fig. 7, which yields  $c_2 = 1.5$  and  $m = 0.6$ . It should be  
344 mentioned that the values of the three constants were obtained from the data related to  
345 naturally worn sediments or grains of similar shape, and further adjustments could be made  
346 for spheres or other shapes of grains once relevant data are available.

347

#### 348 **Case 4: Mobile sediment bed ( $k_s > D$ )**

349 This case appears to be more complicated than the first three cases because bedload grains  
350 and bed materials become interchangeable. In this case,  $k_s$  is usually greater than  $D$  but  
351 cannot be measured directly. Recking et al. (2008) shows that for flows over flat sediment  
352 beds,  $k_s$  increases from  $2.1D$  without bedload to  $5.3D$  with intense bedload transport. It may  
353 further increase significantly in the presence of bedforms such as ripples and dunes (Chien  
354 and Wan, 1999, van Rijn, 1982). Therefore, when fitting Eq. (16) to the data collected under  
355 mobile bed conditions, it is expected that the relationship of  $\chi$ - $\theta$  will deviate from that for the  
356 condition of  $k_s \approx D$ .

357 Due to difficulties encountered in observations under mobile bed conditions, only limited  
358 data have been reported in the literature. Fernandez Luque and van Beek (1976) conducted  
359 their mobile-bed experiments in a closed conduit with both horizontal and inclined channel  
360 beds. To avoid slope effects on the grain velocity, only the data collected under the horizontal  
361 bed condition are analyzed here. Niño et al.'s (1994) mobile-bed experiments involved two  
362 different natural sediments with mean diameters of 15 and 31 mm. When fitting Fernandez  
363 Luque and van Beek's (1976) data [also see Fernandez Luque (1974)] to Eq. (16), we took  $c_1$   
364 = 13 as in case 3, and then found  $c_2(k_s/D)^m = 2$  (Fig. 8). Similarly, with Niño et al.'s (1994)  
365 data, it can be found that  $c_2(k_s/D)^m = 3.3$ . If taking  $c_2 = 1.5$  and  $m = 0.6$ , it can be shown that  
366  $k_s/D$  varies from 1.6 for Fernandez Luque and van Beek's (1976) data to 3.7 for Niño et al.'s

367 (1994) data. Such variations in  $k_s/D$  are consistent with the observation by Recking et al.  
368 (2008).

369

370

## 371 **Comparisons with Previous Formulas**

372 In this section, all the previous formulas with known coefficients are used to predict the  
373 average velocity of solitary bedload grain, and then the predictions are compared with the  
374 measurements (for turbulent flows only) as summarized in Table 2. The critical Shields  
375 number was calculated using Whitehouse et al.'s (2000) formula that represents the Shields  
376 diagram.

377 Fig. 9 shows the comparisons for the case of  $k_s \approx D$ . The predictions are assessed with the  
378 average of error defined as  $|\text{prediction} - \text{measurement}|/\text{measurement}$  (%), and the errors are  
379 summarized in Table 4. It can be seen that the under-predictions are due to the formulas by  
380 Fernandez Luque and van Beek (1976), Fredsoe and Deigaard (1992), Lee and Hsu (1994),  
381 Lee et al. (2000) and Wu (2008), the over-predictions are associated with the formulas by van  
382 Rijn (1984) and Julien and Bounvilay (2013), and the most acceptable predictions are given  
383 by the formula by Hu and Hui (1996) and the present study [i.e. Eq. (16) with  $c_1 = 13.0$ ,  $c_2 =$   
384  $1.5$  and  $m = 0.6$ ]. It is also noted that Eq. (16) differs from the previous formulas except for  
385 the one by Julien and Bounvilay (2013) in that the previous formulas may fail to predict the  
386 grain velocity at low bed shear stresses because of the use of the critical bed shear stress.

387 Among the previous studies, the effect of bed roughness was considered only by Julien  
388 and Bounvilay (2013), their formula is further used to calculate the bedload grain velocity for  
389 rough beds of  $k_s \leq D$ . As shown in Fig. 10(a), the predictions by Julien and Bounvilay's  
390 formula agree very well with their own data for  $k_s < D$ , but showing large deviations for the  
391 other data with  $k_s \approx D$ , which results in an average error of 45.7%. In comparison, Eq. (16)

392 yields consistent predictions, as shown in Fig. 10(b), of which the average error is 15.7%.  
393 This improvement is significant, but not surprising because the three constants were  
394 calibrated with the same database.

395 Next, Eq. (16) is applied to predict the bedload grain velocity for all the data as  
396 summarized in Table 2, in spite of possible effects of the grain shape. The comparison (for  
397 turbulent flows) is presented in Fig. 11, and the average of the prediction errors is 18.2%. It  
398 should be noted that none of the previous formulas can be applied for both smooth and rough  
399 bed conditions, and the formula by Julien and Bounvilay (2013), though involving the bed  
400 roughness height, cannot be applied to smooth beds. The good agreement shown in Fig. 11  
401 implies that the proposed formula, Eq. (16), is generally applicable for the prediction of the  
402 bedload grain velocity under rigid (both smooth and rough) channel bed conditions. The three  
403 constants could be further adjusted to take into account the effect of the grain shape.

404

405

## 406 **Conclusions**

407 By reviewing the previous formulas developed for predicting average velocity of solitary  
408 bedload grain, two observations are made. First, almost all the previous formulas are subject  
409 to the evaluation of the critical condition for incipient sediment motion, and thus may fail to  
410 predict the grain velocity at bed shear stresses lower than the critical values. Second, there is  
411 no formula that is applicable for both smooth and rough bed conditions. To avoid the  
412 shortcomings, a new formula is then developed for the calculation of the bedload grain  
413 velocity in the context of spatial average.

414 The new formula shows that the normalized grain velocity varies with the dimensionless  
415 bed shear stress or the Shields number and also the ratio of bed roughness height to the grain  
416 diameter. This result applies for motion of solitary grains over rough beds, but it may suggest

417 that bedload transport rate not only depends on the Shields number but also the equivalent  
418 bed roughness height even under mobile bed conditions and such dependence may become  
419 pronounced for non-uniform sediments.

420 To calibrate the proposed formula, the experimental data from the present study and those  
421 reported in the literature are compiled and analyzed under different flow conditions,  
422 including laminar flows over a smooth bed, turbulent flows over a smooth bed, turbulent  
423 flows over a rigid rough bed with varying roughness heights, and turbulent flow over a  
424 mobile bed. In comparison with seven previous formulas, the new formula, Eq. (16),  
425 performs the best in the prediction of the bedload grain velocity for both smooth and rough  
426 configurations. In addition, Eq. (16) could be also applied for the mobile bed condition  
427 provided that the equivalent roughness height of the sediment bed is available. However, it  
428 should be noted that the present formula was calibrated using the data largely for naturally-  
429 worn coarse grains. Further studies should be conducted to investigate viscous effect on the  
430 average velocity of fine bedload grains.

431

432

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436

### 437 **Notation**

438 The following symbols are used in this paper:

439  $C_D$  drag coefficient;

440  $c_1$  coefficient;

441  $c_2$  coefficient;

442  $c_3$  coefficient;

443	$D$	grain diameter;
444	$D^*$	dimensionless grain diameter;
445	$F_D$	drag force;
446	$F_L$	lift force;
447	$F_R$	bed resistance force;
448	$g$	gravitational acceleration;
449	$h$	flow depth;
450	$k_s$	equivalent roughness height of the channel bed;
451	$k_s^+$	$= u_* k_s / \nu$ ;
452	$m$	exponent;
453	$Re_*$	$= u_* D / \nu$ ;
454	$S$	channel slope;
455	$s$	distance from the center of moving grain to boundary;
456	$T$	temperature;
457	$T^*$	transport stage parameter;
458	$u_*$	shear velocity;
459	$u_b$	unobstructed, near-bed flow velocity at $y_b$ ;
460	$u_s$	average grain velocity;
461	$W$	submerged grain weight;
462	$y_b$	representative distance from the center of moving grain to channel bed;
463	$y_b^+$	$= u_* y_b / \nu$ ;
464	$\alpha$	friction angle;
465	$\beta$	slope angle;
466	$\Delta$	effective specific density [ $= (\rho_s - \rho) / \rho$ ];
467	$\theta$	dimensionless bed shear stress or Shields number;
468	$\theta_c$	critical Shields number for incipient sediment motion;
469	$\nu$	kinematic viscosity of fluid;
470	$\rho$	fluid density;
471	$\rho_s$	grain density; and
472	$\chi$	dimensionless grain velocity.

473

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**Table 1.** Previous Formulas for Calculating Bedload Grain Velocity

Investigator	$\chi$	Note
Einstein (1950)	$c\theta^{0.5}$	$c = \text{constant}$ ; $\theta'$ is grain-related Shields number
Bagnold (1973)	$U/(\Delta gD)^{0.5} - 5.75\theta^{0.5}\log[0.37h/(nD)] - w/(\Delta gD)^{0.5}$	$n = 1.4(\theta/\theta_c)^{0.6}$ ; $h = \text{flow depth}$ ; $U = \text{depth-averaged flow velocity}$ ; $w = \text{settling velocity}$
Fernandez Luque and van Beek (1976)	$11.5(\theta^{0.5} - 0.7\theta_c^{0.5})$	
Abbott and Francis (1977)	$c(\theta^{0.5} - 0.14)$	$c = 13.5-14.3$
van Rijn (1984)	$(9+2.6\log D_*)\theta^{0.5} - 8\theta_c^{0.5}$	
Bridge and Dominic (1984)	$c(\theta^{0.5} - \theta_c^{0.5})$	$c = 6.6-15.6$
Fredsoe and Deigaard (1992)	$10(\theta^{0.5} - 0.7\theta_c^{0.5})$	
Hu and Hui (1996)	$19.5\theta^{0.5} - 5.1\theta_c^{0.5}$ (for smooth bed) $11.9\theta^{0.5} - 5.2\theta_c^{0.5}$ (for rough bed)	
Lee and Hsu (1994)	$11.5\theta^{0.5}(\theta - \theta_c)^{0.174}$	
Lee et al. (2000)	$3.14D_*^{0.17}(\theta/\theta_c - 1)^{0.18}\theta^{0.5}$	
Wu (2008)	$1.64(\theta/\theta_c - 1)^{0.5}$	
Julien and Bounvilay (2013)	$11.5\theta^{0.95}D_*^{0.21}(D/k_s)^{0.36}\Delta^{-0.28}$	

**Table 2.** Summary of Experimental Data Used for Analysis

Investigator	Bed Roughness height $k_s$ (mm)	Tested grain			Shields number $\theta$	Dimensionless grain diameter $D_*$	No. of observations
		Shape	Diameter D (mm)	Specific gravity $\rho_s/\rho$			
(1) Smooth bed ( $k_s \ll D$ )							
Parsons (1972)		spherical	0.25-0.85		0.0037-0.041	7-26	31
Hu and Hui (1996)		spherical; angular	2.3-3.6	1.04-2.65	0.15-1.41	16.8-88.5	10
Julien and Bounvilay (2013)		spherical; angular	1.57-29.3	2.6-8.02	0.000076-0.016	62-778	172
Present study		spherical	11	1.14-2.59	0.0019-0.16	136-306	52
(2) Rigid sediment bed ( $k_s \approx D$ )							
Francis (1973)	7.5	naturally-worn	7.5	1.04-2.42	0.042-2.41	55-180	30
Abbott and Francis (1977)	7.2	naturally-worn	8.82	1.20-2.86	0.035-1.12	110-232	68
Lee and Hsu (1994)	1.36	naturally-worn	1.36	2.64	0.059-0.50	34-62	9
Nina and Garcia (1998)	0.5	naturally-worn	0.5	2.65	0.053-0.127	13	6
Lee et al. (2000)	6		6	1.08-1.38	0.066-0.55	55-93	6

Julien and Bounvilay (2013)	2.4-3.4	naturally worn	2.4-3.4	2.65	0.035-0.067	61-86	6
Present study	11	spherical	11	1.14-2.59	0.012-0.16	136-306	44
(3) Rigid sediment bed ( $k_s < D$ )							
Julien and Bounvilay (2013)	1.2-3.4	spherical; angular	2.4-29.3	2.6-8.02	0.00079-0.090	61-797	351
Lee and Hsu (1994)	1.36	Naturally-worn	2.47	2.64	0.142-0.199	62	3
(4) Mobile sediment bed ( $k_s > D$ )							
Fernandez Luque and van Beek (1974)		naturally-worn	0.9-3.3	1.34-4.58	0.044-0.076	22-83	17
Nino and Garcia (1994)		naturally-worn	15-31	2.65	0.084-0.15	379-784	6

**Table 3.** Summary of Experimental Data Collected in the Present Study (sphere diameter = 11 mm, and channel width = 100 mm).

Test No.	Bed roughness height $k_s$ (m)	$\Delta$ ( $= \rho_s/\rho - 1$ )	Flow depth h (m)	Slope	Shear velocity $u_*$ (m/s)	Grain velocity $u_s$ (m/s)
1	0.011	0.14	0.015	0.005	0.028	0.196
2	0.011	0.14	0.015	0.007	0.032	0.230
3	0.011	0.14	0.015	0.009	0.037	0.277
4	0.011	0.14	0.015	0.011	0.040	0.321
5	0.011	0.14	0.015	0.012	0.042	0.350
6	0.011	0.14	0.015	0.014	0.045	0.373
7	0.011	0.14	0.015	0.016	0.049	0.412
8	0.011	0.48	0.0165	0.008	0.037	0.181
9	0.011	0.48	0.018	0.008	0.038	0.191
10	0.011	0.48	0.0195	0.008	0.040	0.201
11	0.011	0.48	0.0215	0.008	0.042	0.215
12	0.011	0.48	0.014	0.015	0.045	0.272
13	0.011	0.48	0.015	0.015	0.047	0.293
14	0.011	0.48	0.016	0.015	0.048	0.303
15	0.011	0.48	0.0195	0.015	0.054	0.351
16	0.011	0.48	0.013	0.019	0.049	0.327
17	0.011	0.48	0.014	0.019	0.051	0.340
18	0.011	0.48	0.015	0.019	0.053	0.369
19	0.011	0.48	0.017	0.019	0.056	0.398
20	0.011	0.48	0.015	0.004	0.024	0.078
21	0.011	0.48	0.015	0.006	0.030	0.144
22	0.011	0.48	0.015	0.007	0.033	0.172
23	0.011	0.48	0.015	0.011	0.041	0.269
24	0.011	0.48	0.015	0.016	0.048	0.337
25	0.011	0.48	0.015	0.020	0.054	0.388
26	0.011	0.48	0.015	0.021	0.056	0.409
27	0.011	0.48	0.015	0.024	0.060	0.443
28	0.011	0.48	0.015	0.027	0.063	0.490
29	0.011	0.48	0.015	0.029	0.065	0.488
30	0.011	1.59	0.014	0.029	0.063	0.233
31	0.011	1.59	0.0145	0.031	0.066	0.302
32	0.011	1.59	0.014	0.031	0.065	0.281
33	0.011	1.59	0.0155	0.024	0.060	0.203
34	0.011	1.59	0.015	0.024	0.059	0.185
35	0.011	1.59	0.015	0.015	0.047	0.114
36	0.011	1.59	0.0165	0.015	0.049	0.137
37	0.011	1.59	0.015	0.019	0.053	0.202
38	0.011	1.59	0.0165	0.019	0.055	0.215

39	0.011	1.59	0.015	0.022	0.057	0.241
40	0.011	1.59	0.0165	0.022	0.060	0.279
41	0.011	1.59	0.015	0.025	0.061	0.306
42	0.011	1.59	0.0165	0.025	0.064	0.318
43	0.011	1.59	0.015	0.028	0.064	0.346
44	0.011	1.59	0.0165	0.028	0.067	0.389
45	0	1.59	0.013	0.003	0.0181	0.229
46	0	1.59	0.016	0.003	0.0201	0.251
47	0	1.59	0.019	0.003	0.0219	0.265
48	0	1.59	0.022	0.003	0.0236	0.290
49	0	1.59	0.013	0.005	0.0258	0.361
50	0	1.59	0.0155	0.005	0.0282	0.387
51	0	1.59	0.018	0.005	0.0304	0.415
52	0	1.59	0.013	0.010	0.0354	0.448
53	0	1.59	0.0165	0.010	0.0399	0.481
54	0	1.59	0.018	0.010	0.0417	0.502
55	0	1.59	0.0125	0.013	0.0400	0.521
56	0	1.59	0.014	0.013	0.0423	0.531
57	0	1.59	0.016	0.013	0.0452	0.558
58	0	1.59	0.0175	0.013	0.0473	0.568
59	0	1.59	0.013	0.015	0.0441	0.591
60	0	1.59	0.014	0.015	0.0457	0.614
61	0	1.59	0.016	0.015	0.0489	0.633
62	0	1.59	0.0165	0.015	0.0497	0.639
63	0	0.48	0.013	0.003	0.0181	0.242
64	0	0.48	0.016	0.003	0.0201	0.271
65	0	0.48	0.019	0.003	0.0219	0.289
66	0	0.48	0.022	0.003	0.0236	0.308
67	0	0.48	0.013	0.005	0.0258	0.373
68	0	0.48	0.0155	0.005	0.0282	0.408
69	0	0.48	0.018	0.005	0.0304	0.453
70	0	0.48	0.013	0.010	0.0354	0.457
71	0	0.48	0.0165	0.010	0.0399	0.513
72	0	0.48	0.018	0.010	0.0417	0.547
73	0	0.48	0.0125	0.013	0.0400	0.533
74	0	0.48	0.014	0.013	0.0423	0.547
75	0	0.48	0.016	0.013	0.0452	0.592
76	0	0.48	0.0175	0.013	0.0473	0.613
77	0	0.48	0.0165	0.015	0.0497	0.684
78	0	0.48	0.0155	0.015	0.0481	0.673
79	0	0.48	0.014	0.015	0.0457	0.668
80	0	0.14	0.013	0.003	0.0181	0.268
81	0	0.14	0.016	0.003	0.0201	0.304
82	0	0.14	0.019	0.003	0.0219	0.328

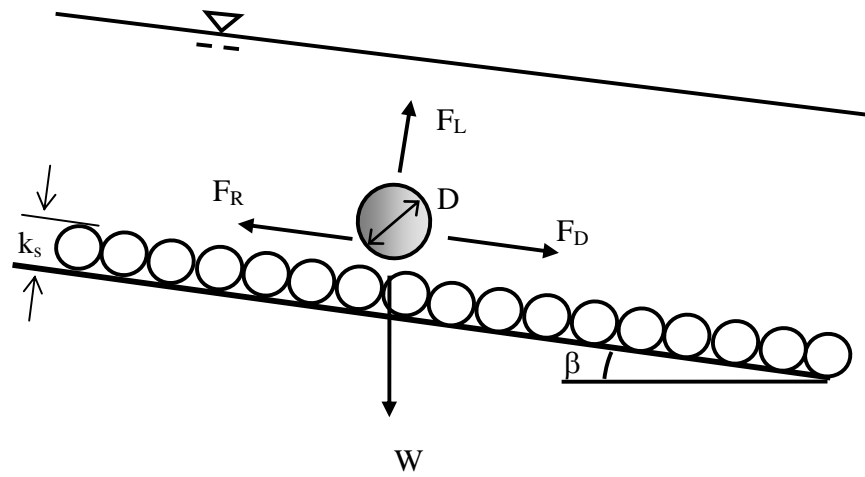
83	0	0.14	0.022	0.003	0.0236	0.351
84	0	0.14	0.013	0.005	0.0258	0.391
85	0	0.14	0.0155	0.005	0.0282	0.441
86	0	0.14	0.018	0.005	0.030	0.496
87	0	0.14	0.013	0.010	0.035	0.495
88	0	0.14	0.0165	0.010	0.040	0.556
89	0	0.14	0.018	0.010	0.042	0.596
90	0	0.14	0.0125	0.013	0.040	0.577
91	0	0.14	0.014	0.013	0.042	0.599
92	0	0.14	0.016	0.013	0.045	0.640
93	0	0.14	0.0175	0.013	0.047	0.666
94	0	0.14	0.0165	0.015	0.050	0.747
95	0	0.14	0.0155	0.015	0.048	0.738
96	0	0.14	0.014	0.015	0.046	0.711

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**Table 4.** Summary of Prediction Errors

Investigator	Average error (%)
Rough beds with $k_s = D$	
Fernandez Luque and van Beek (1976)	23.1
van Rijn (1984)	14.2
Fredsoe and Deigaard (1992)	32.0
Lee and Hsu (1994)	16.2
Hu and Hui (1996)	11.0
Lee et al. (2000)	13.9
Wu (2008)	37.1
Julien and Bounvilay (2013)	145.3
Present study	9.9
Rough beds with $k_s \leq D$	
Julien and Bounvilay (2013)	45.7
Present study	15.7
Smooth and rough beds	
Present study	18.2

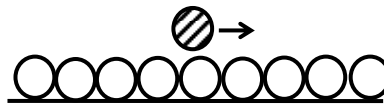




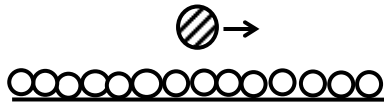
**Figure 1.** Forces exerting on a solitary bedload grain.



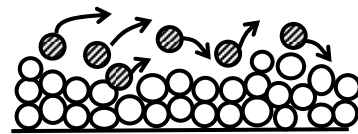
(a) Smooth bed ( $k_s \ll D$ )



(b) Rigid sediment bed ( $k_s = D$ )

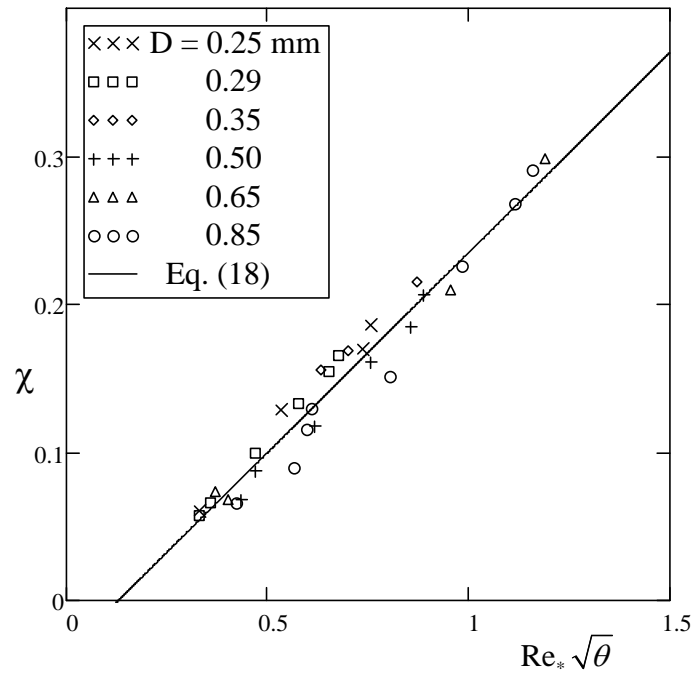


(c) Rigid sediment bed ( $k_s < D$ )

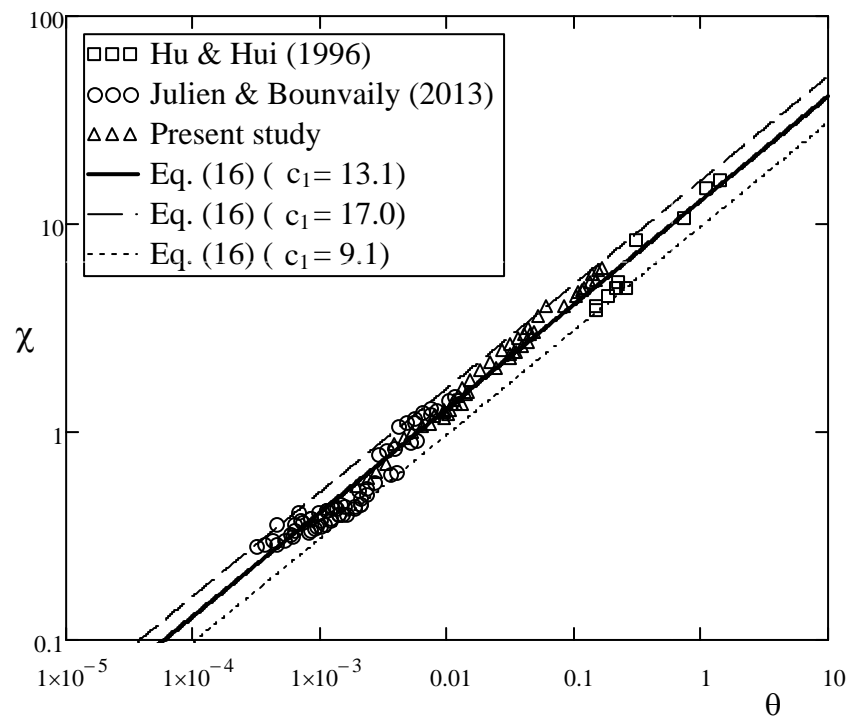


(d) Mobile sediment bed ( $k_s > D$ )

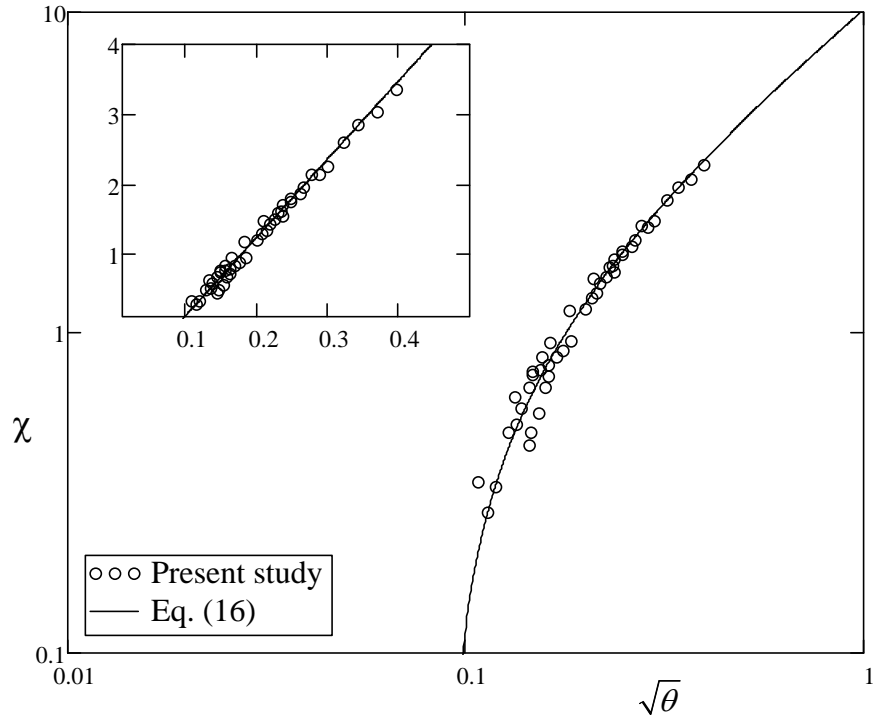
**Figure 2.** Bedload grains moving over four different channel beds.



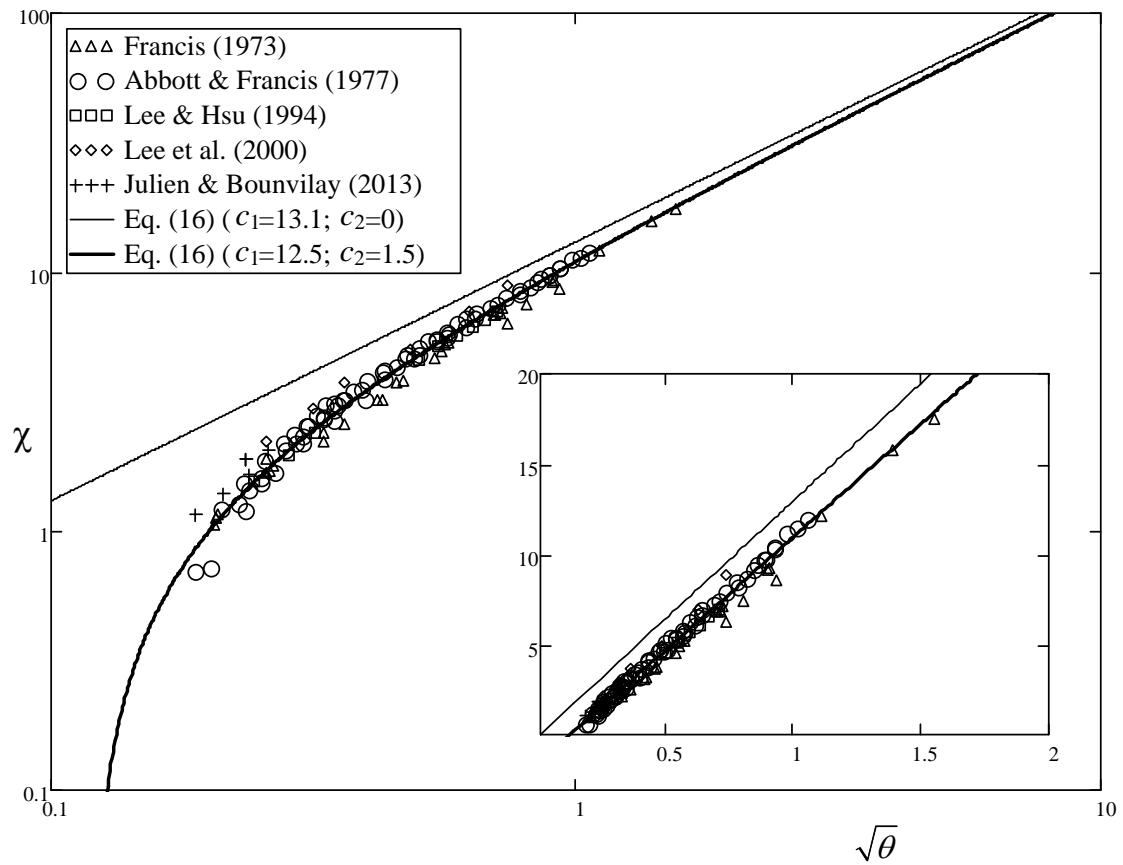
**Figure 3.** Variation of dimensionless bedload grain velocity for laminar flows over a smooth bed. The experiments were conducted with glass beads (*Parsons 1972*). Eq. (18) was plotted by taking  $c_3 = 0.27$  and the last term as a constant ( $= 0.03$ ).



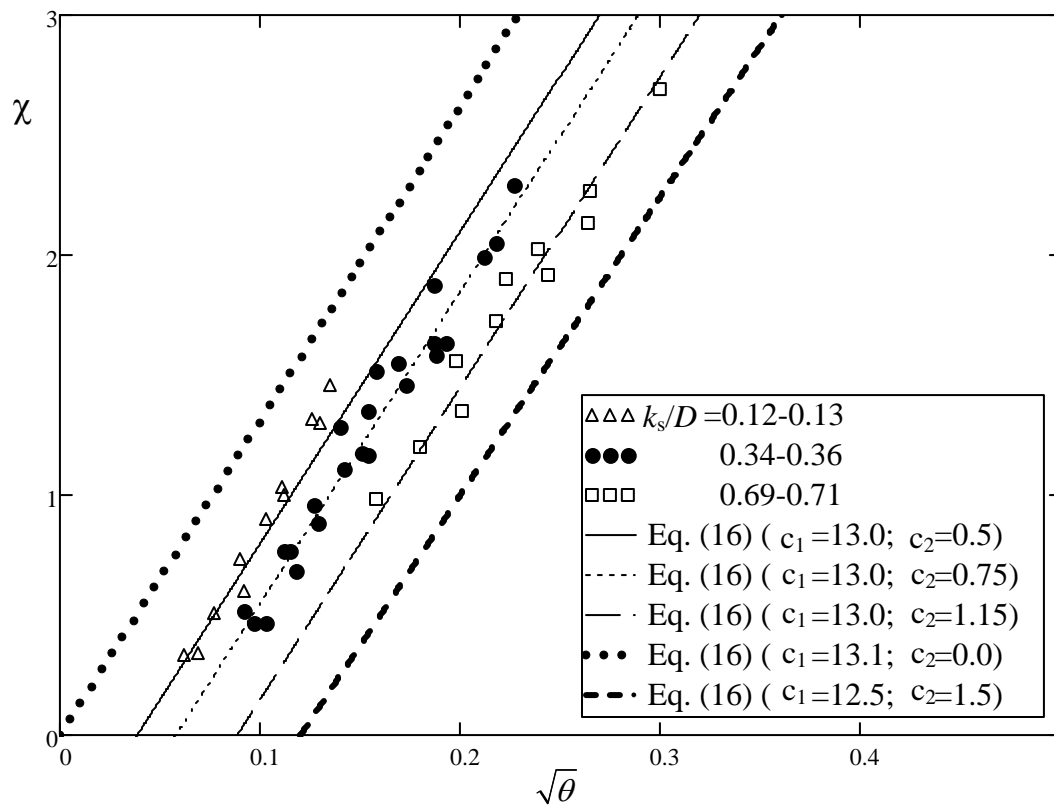
**Figure 4.** Variation of dimensionless grain velocity for turbulent flows over a smooth bed. Eq. (16) was used to fit the data with  $c_1 = 13.1$  (with a variation of  $\pm 30\%$ ) and  $c_2 = 0$ .



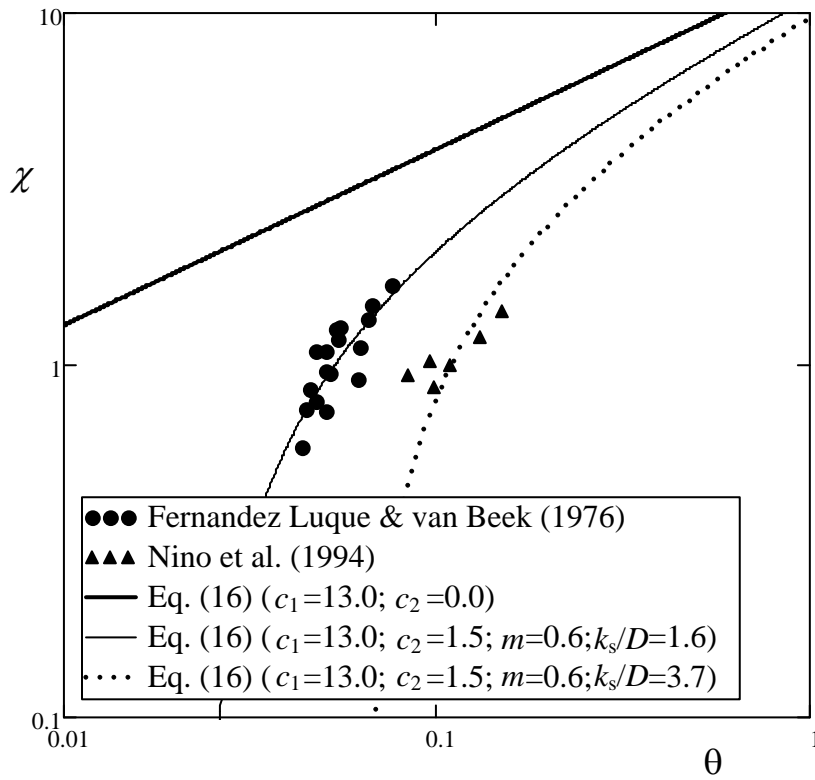
**Figure 5.** Variation of dimensionless grain velocity for spheres with  $k_s = D$ . The data were collected in the present study and fitted with Eq. (16) by taking  $c_1 = 12.1$  and  $c_2 = 1.0$ . The insert shows the same data plotted with linear coordinates.



**Figure 6.** Variation of dimensionless grain velocity for rough beds comprising naturally-worn grains with  $k_s = D$ . Eq. (16) was fitted to the data by taking  $c_1 = 12.5$  and  $c_2 = 1.5$ . The insert shows the same data plotted with linear coordinates.

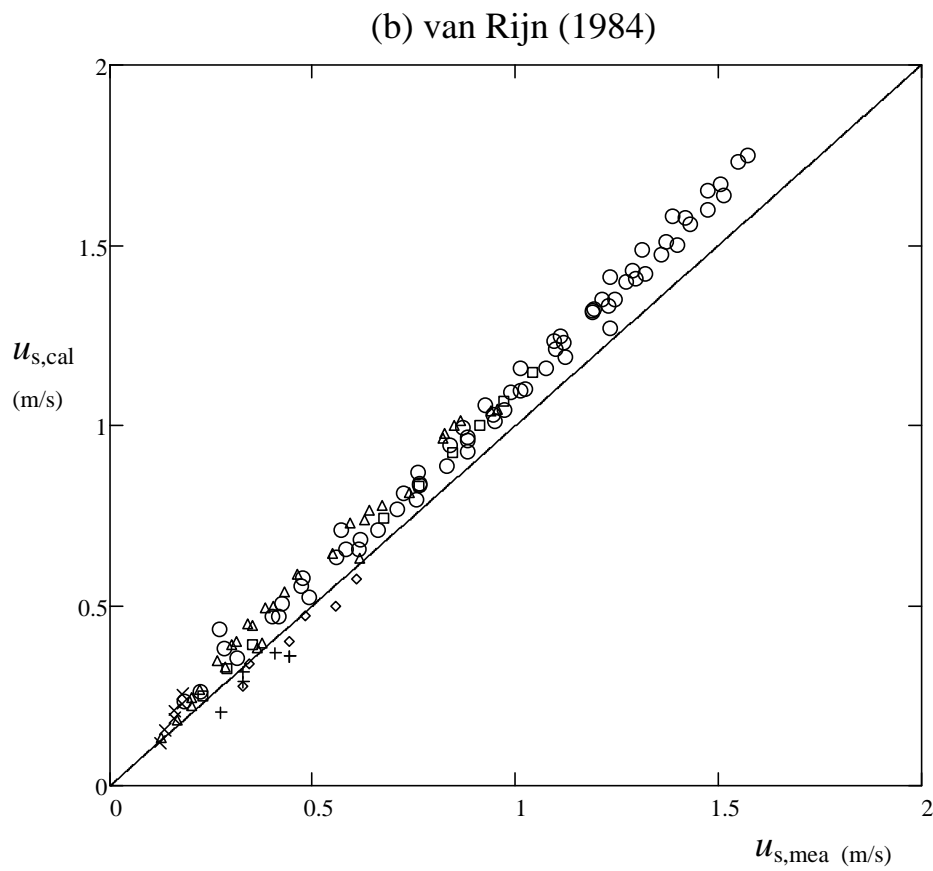
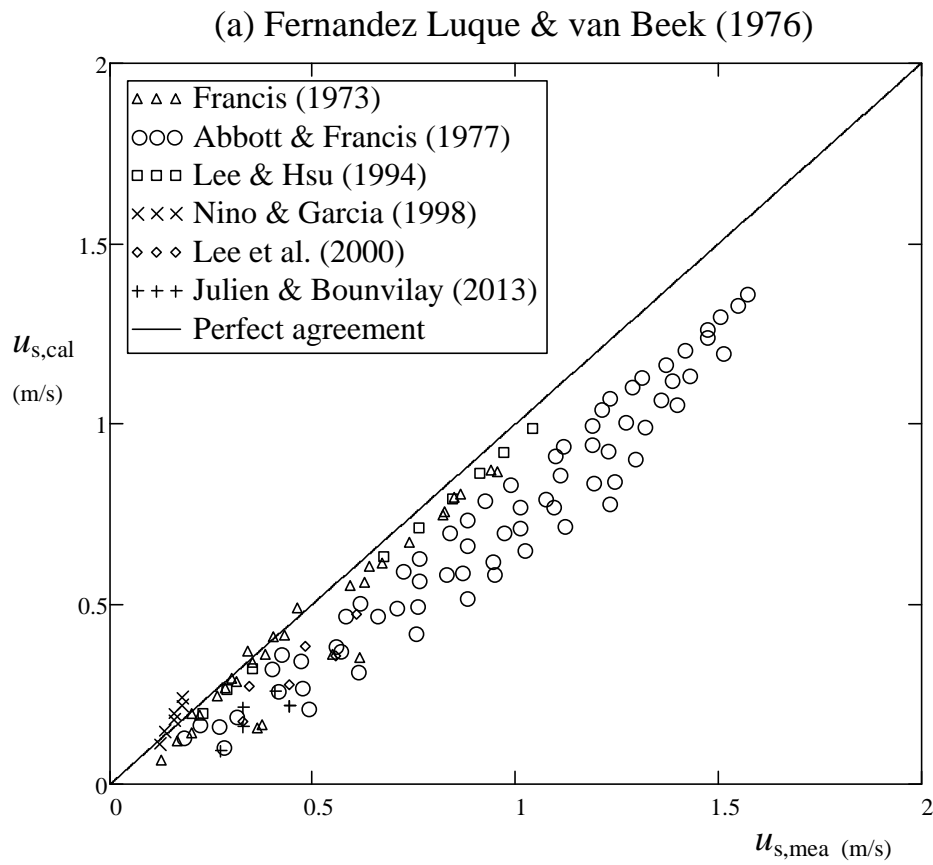


**Figure 7.** Effects of  $k_s/D$  on the relationship of  $\chi-\sqrt{\theta}$ . The data were reported by Julien and Bounvilay (2013). Eq. (16) was superimposed with different coefficients. The uppermost line represents the data for the smooth bed while the lowest line represents the data for  $k_s = D$ .

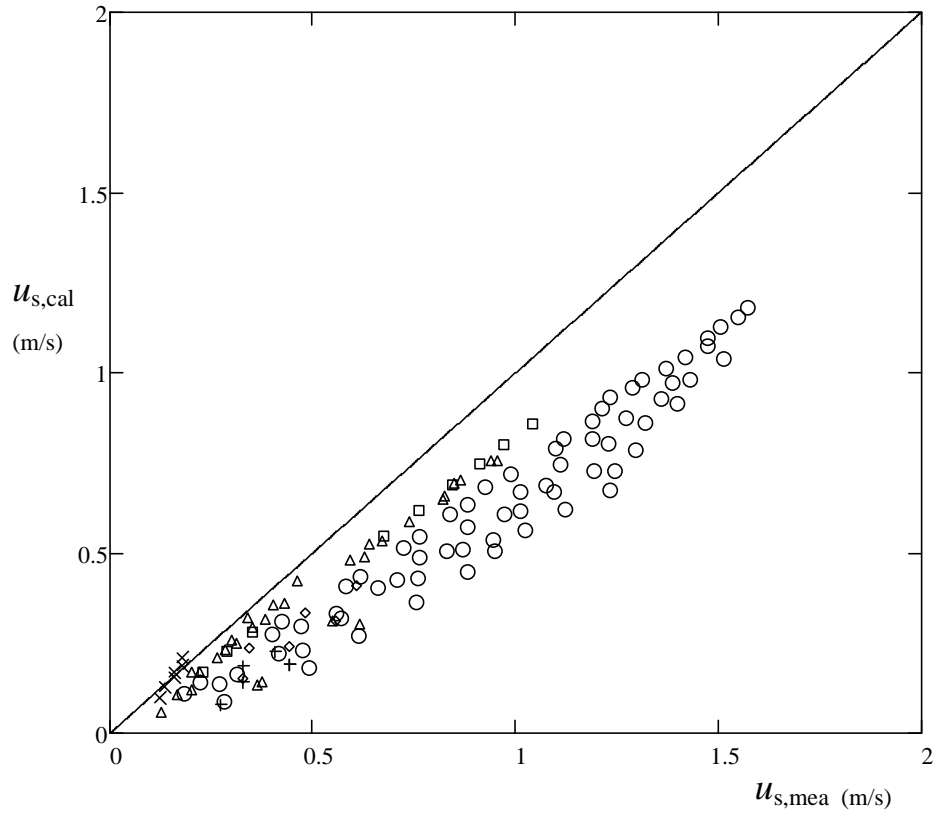


**Figure 8.** Fitting of Eq. (16) to experimental data for the mobile bed conditions.

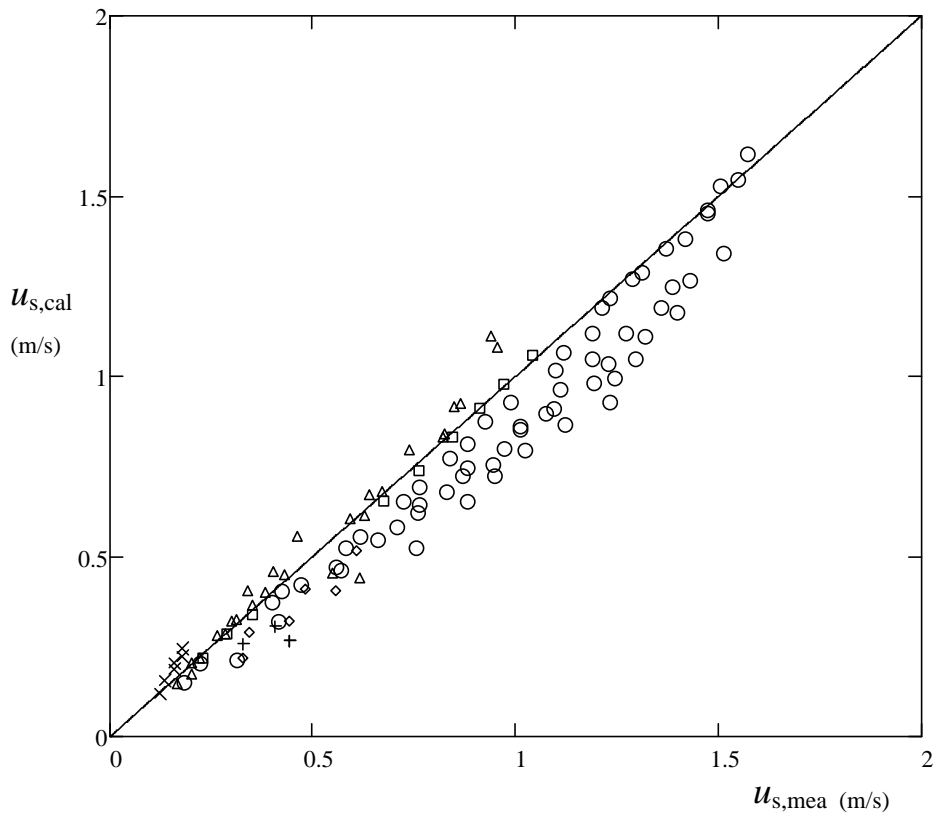




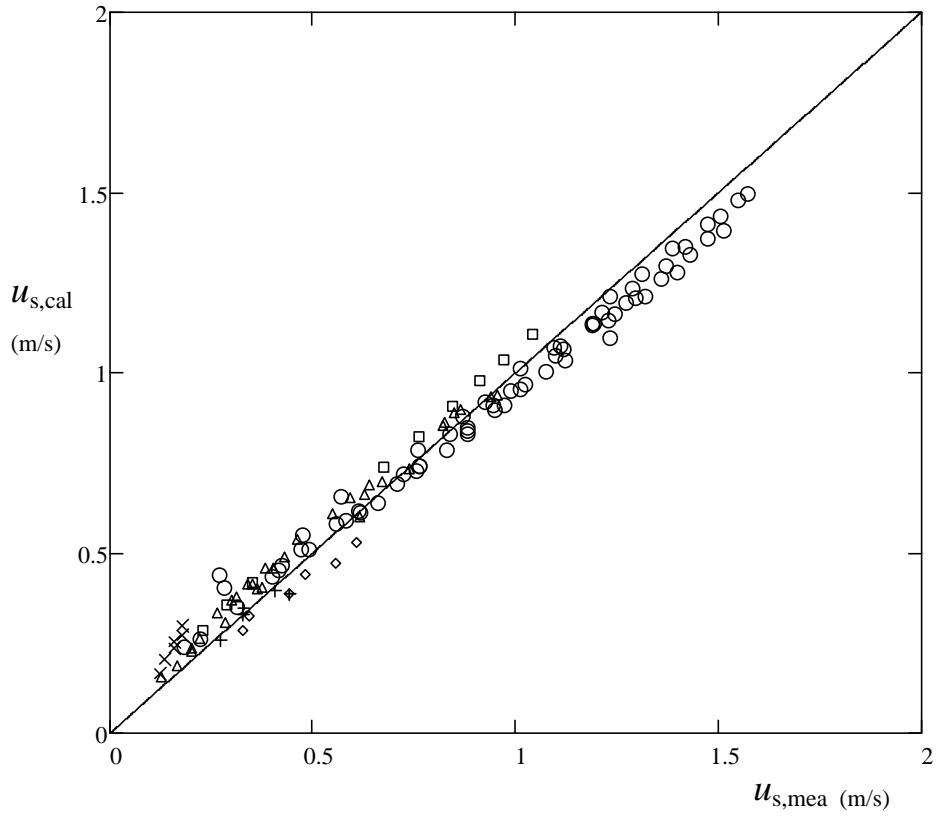
(c) Fredsoe & Deigaard (1992)



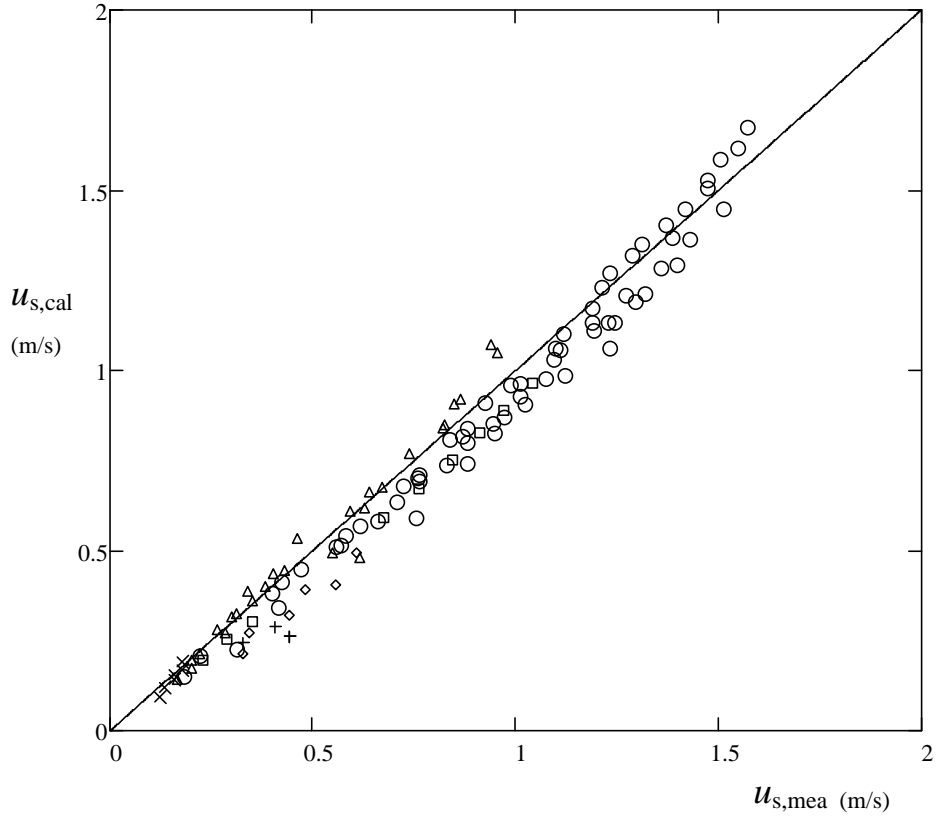
(d) Lee & Hsu (1994)



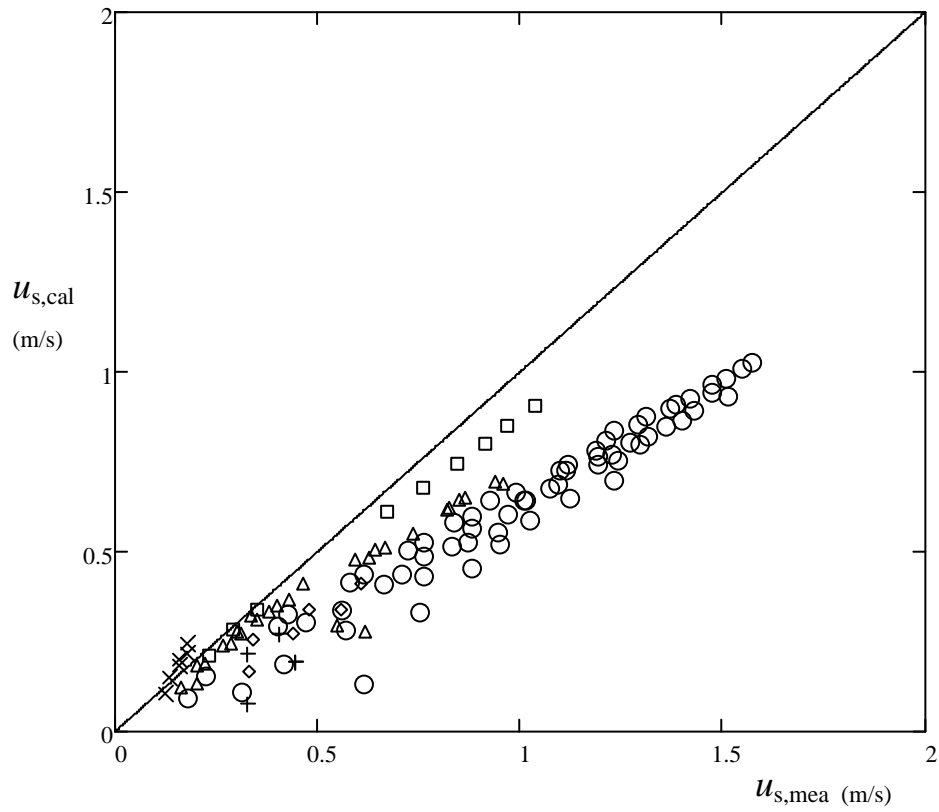
(e) Hu & Hui (1996)



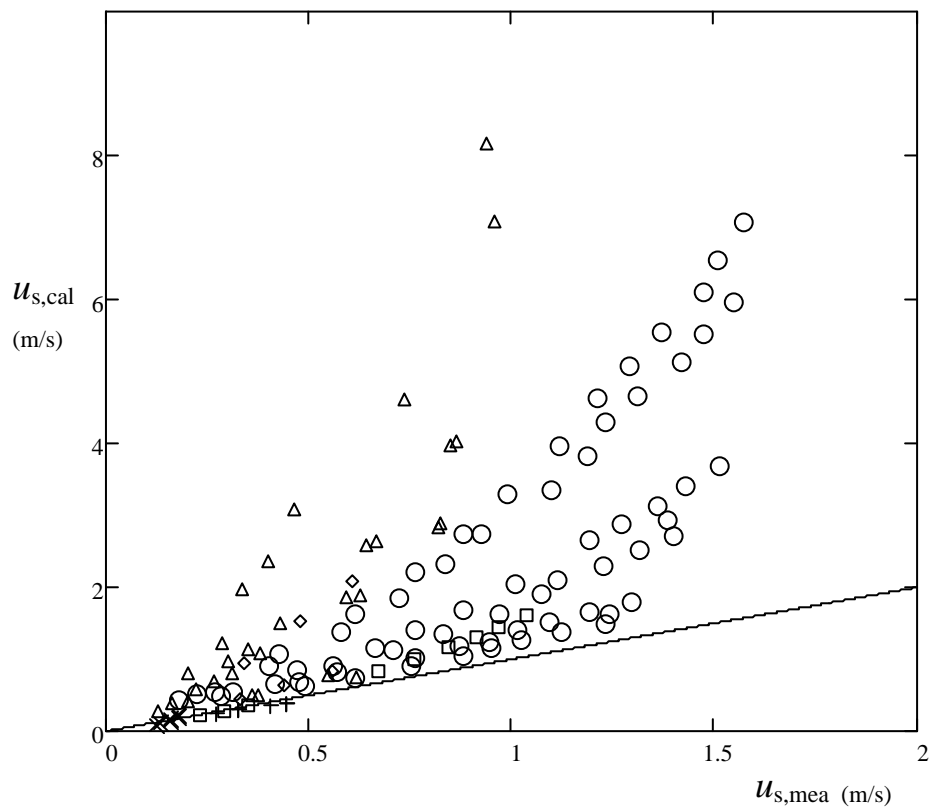
(f) Lee et al. (2000)

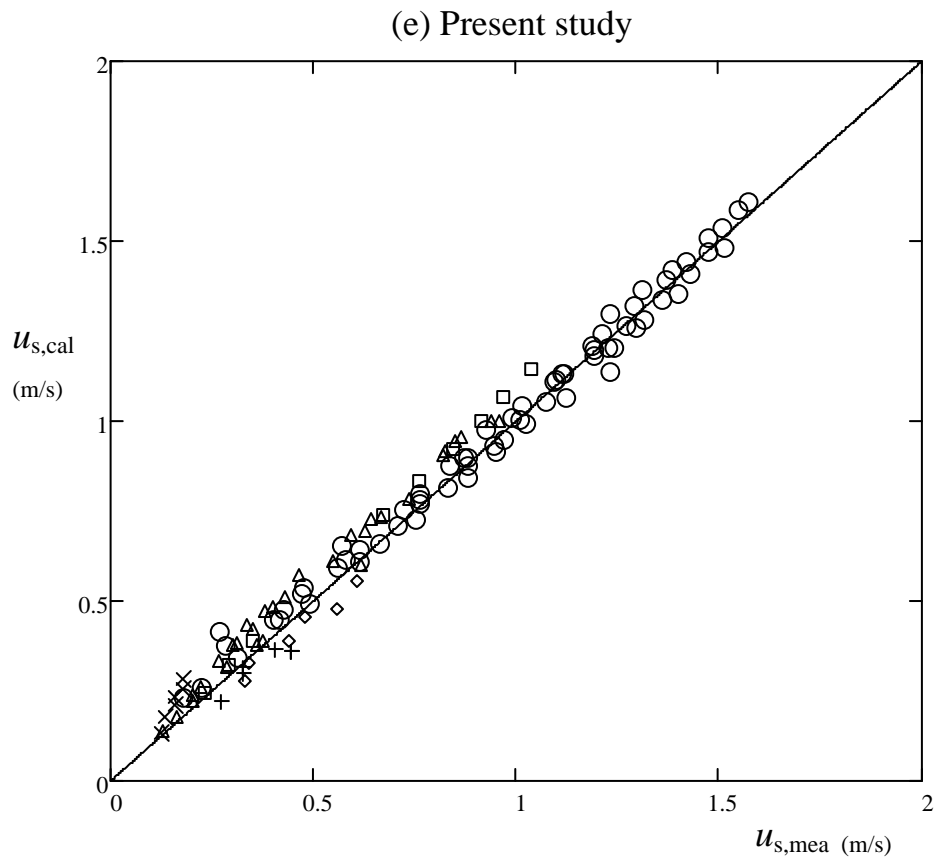


(g) Wu (2008)

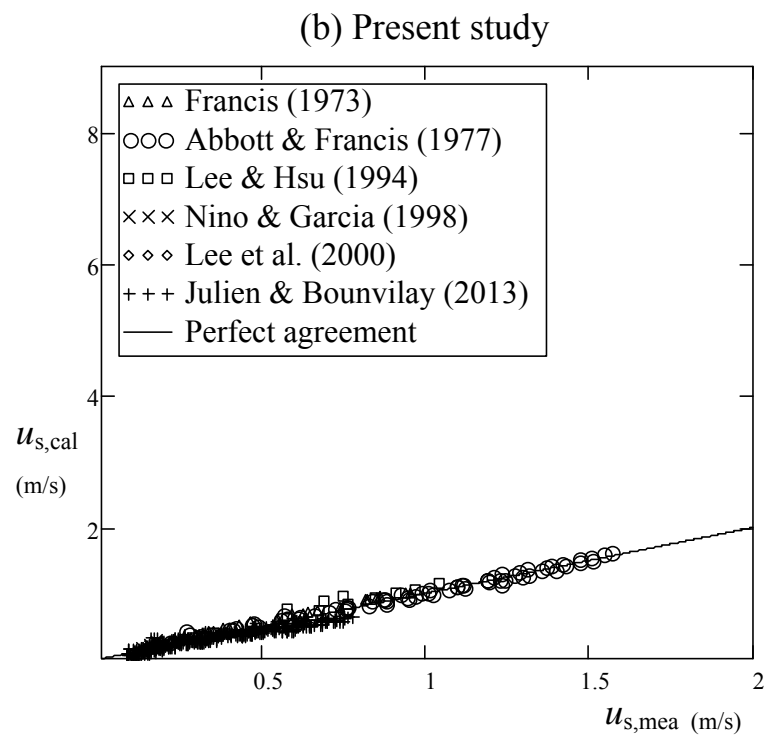
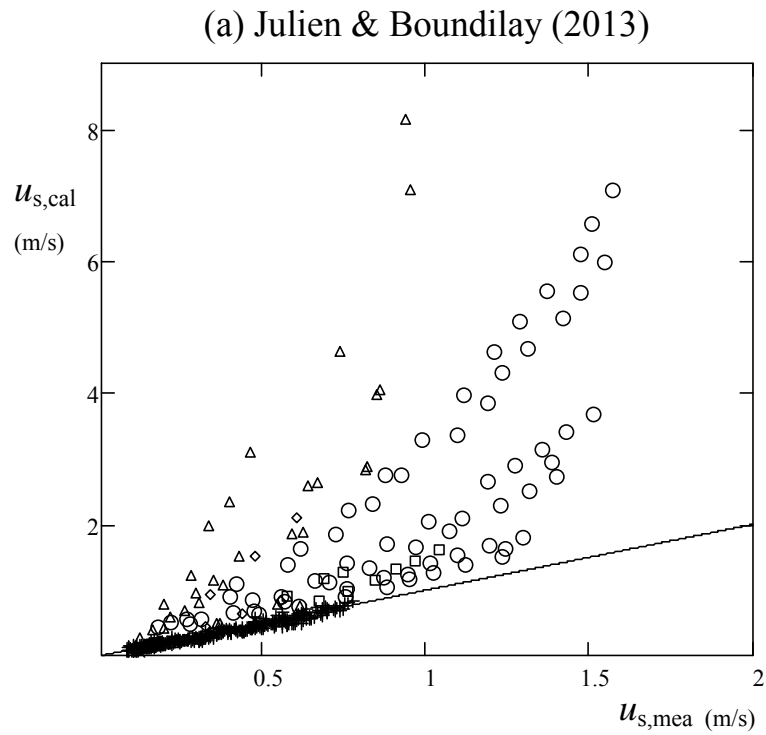


(h) Julien & Boundilay (2013)

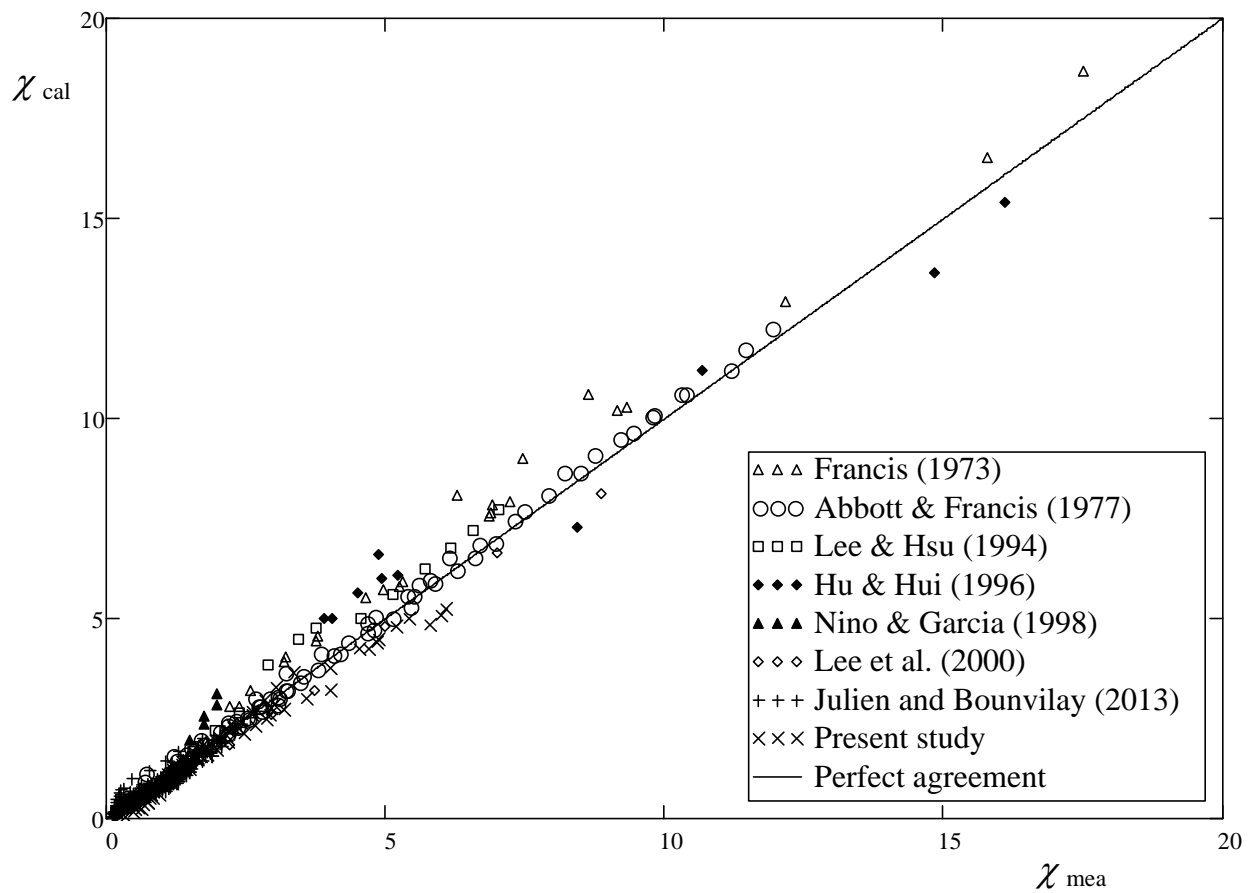




**Figure 9.** Comparison of calculated and measured bedload grain velocities for rough beds comprising naturally-worn grains with  $k_s = D$ . The same symbols are used for figures (a) to (h).



**Figure 10.** Comparison of calculated and measured bedload grain velocities for rough beds comprising naturally-worn grains with  $k_s \leq D$ . The same symbols are used for figures (a) and (b).



**Figure 11.** Comparison of calculated and measured dimensionless bedload grain velocities for both smooth and rough bed conditions. No differentiation was made between spherical and non-spherical grains. Eq. (16) with  $c_1 = 13$ ,  $c_2 = 1.5$  and  $m = 0.6$  was used for calculations.

## **Figure Captions**

**Fig. 1.** Forces exerting on a solitary bedload grain.

**Fig. 2.** Bedload grains moving over four different channel beds.

**Fig. 3.** Variation of dimensionless bedload grain velocity for laminar flows over a smooth bed. The experiments were conducted with glass beads (Parsons 1972). Eq. (18) was plotted by taking  $c_3 = 0.27$  and the last term as a constant ( $= 0.03$ ).

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**Fig. 5.** Variation of dimensionless grain velocity for spheres with  $k_s = D$ . The data were collected in the present study and fitted with Eq. (16) by taking  $c_1 = 11.2$  and  $c_2 = 1.0$ . The insert shows the same data plotted with linear coordinates.

**Fig. 6.** Variation of dimensionless grain velocity for rough beds comprising naturally-worn grains with  $k_s \approx D$ . Eq. (16) was used to fit the data with  $c_1 = 12.5$  and  $c_2 = 1.5$ . The insert shows the same data plotted with linear coordinates.

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**Fig. 8.** Fitting of Eq. (16) to experimental data for the mobile bed conditions.

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